

Technical Document 1074 March 1987

Estimate of Electro-Optical Meteorological Parameters from Satellite-Detected Radiances

> T. H. Vonder Haar M. Wetzel

> > METSAT, Inc.



Approved for public release, distribution is unlimited

The views and conclusions contained in this report are those of the authors and should not be interpreted as representing the official policies. either expressed or implied, of the Naval Ocean Systems Center or the U.S. Government

NAVAL OCEAN SYSTEMS CENTER

San Diego, California 92152-5000

E. G. SCHWEIZER, CAPT, USN
Commander

R. M. HILLYER
Technical Director

ADMINISTRATIVE INFORMATION

This work was performed for the Office of Naval Technology, Office of the Chief of Naval Research, Arlington, VA 22217, under program element 62759N. Contract N66001-85-C-0202 was performed by METSAT, Inc., 515 South Howes, Fort Collins, CO 80521.

Released by H.V. Hitney, Head Tropospheric Branch Under authority of J.H. Richter, Head Ocean and Atmospheric Sciences Division

	R	EPORT DOCUME	NTATION PAG	3E			
1a REPORT SECURITY CLASSIFICATION		1	16 RESTRICTIVE MARKINGS				
UNCLASSIFIEI)							
2. SECURITY CLASSIFICATION AUTHORITY			3 DISTRIBUTION AVAILABI		- 11 - 11 - 1		
26 DECLASSIFICATION DOWNGRADING SCHED	ULE		Approved for pu	DHC release; QI	atribution is un	limited.	
4 PERFORMING ORGANIZATION REPORT NUMBER	ER(S)		5 MONITORING ORGANIZAT	TION REPORT NUMBER	S		
			NOSC TD 1074				
6a NAME OF PERFORMING ORGANIZATION		6b OFFICE SYMBOL (if applicable)	78 NAME OF MONITORING	ORGANIZATION			
METSAT. Inc.			Naval Ocean Sys	tems Center			
6. ADDRESS (City State and ZIP Code.			76 ADDRESS (City, State and	1 ZIP Codei			
515 South Howes Fort Collins. Co 80521			Tropospheric Bra San Diego, CA				
8a NAME OF FUNDING SPONSORING ORGANIZA	ATION	8b OFFICE SYMBOL (if applicable)	9 PROCUREMENT INSTRUM	ENT IDENTIFICATION N	UMBER		
Office of Naval Technology		ONT	N66001-85-C-0202	1			
Bc ADDRESS (City State and ZIP Code)		<u> </u>	10 SOURCE OF FUNDING N	UMBERS	- <u> </u>	· · · · · · · · · · · · · · · · · · ·	
			PROGRAM ELEMENT NO	PROJECT NO	TASK NO	AGENCY ACCESSION NO	
Office of the Chief of Naval Arlington, VA 22217	Research		62759N	W 59551	540-SXB3	DN888 715	
11 TITLE Unclude Security Classification;					940-3XB3	7 2.4000 .10	
Estimate of Electro-Optical M 12 PERSONAL AUTHORIS		Parameters from Sate	llite-Detected Radi	ances			
T.H. Vonder Haar and M. W	etzel				-		
13a TYPE OF REPORT	13b TIME COVER	ED g 1985 to Aug 1986	14 DATE OF REPORT (Year,	Month, Dayi	15 PAGE COL	INT	
Final	FROM AU	10 Aug 1300	March 1987		59		
17 COSATI CODES		18 SUBJECT TERMS (Commune of Electro-optics	in reverse if necessary and idea	wify by block number. propagation			
FIELD GROUP	SUB GROUP	electromagnetic	atmosphere				
		aerosols advanced very high		albedo			
19 ABSTRACT Continue on reverse if necessary a	and identify by black to	<u> </u>	- Testration radiom	- (Avincie)	<u></u>		
This document describes analysis of digital image data model for marine stratus c applications. Summaries of the estimating clear-sky atmospher sensing applications for bound	from GOES, louds, and i he research p ic transmitta	DMSP, and NOAA in the design and testing roject and results are note from satellite data	satellites, developm ing of a spectral provided in this i	ent of a mono radiometer i report. Section	ochromatic radi for marine bo n 2 describes t	ative transfer undary layer he analysis of	
20 DISTRIBUTION AVAILABILITY OF ABSTRACT	_		21 ABSTRACT SECURITY			··	
UNCLASSIFIED UNLIMITED	SAME AS RPT	DTIC USERS	UNCLASSIFIE		I a ore a second		
H Hughes			(619) 225-6520	vwa code	Code 543		

DD FORM 1473, 84 JAN

CONTENTS

1.0	Summary	1
	Satellite Observations of Clear-Sky Radiance	
	2.1 Sensitivity of LOWTRAN 6 to Boundary Layer Conditions	2
	2.2 Correspondence of Satellite and LOWTRAN 6 Radiance	5
	2.3 Image Analysis	9
3.0	Satellite Observations of Marine Boundary Layer Clouds	10
4.0	Summary and Recommendations	13
5.0	Acknowledgements	14
6.0	References	15
Table	25	17
Figure	es	27

1.0 SUMMARY

This report describes work completed during the contract period 15 August 1985 - 14 August 1986. The research tasks include the implementation of atmospheric transmission models, collection and analysis of digital image data from the GOES, DMSP, and NOAA satellites, development of a monochromatic radiative transfer model for marine stratus clouds, and the design and testing of a spectral radiometer for marine boundary layer applications. Two field programs have been carried out in collaboration with NOSC in the San Diego area, and data from previous research studies have been incorporated into this research.

Results from portions of the work were published, and these papers were presented at conferences sponsored by the American Meteorological Society and DoD. Summaries of the research project and results are provided in the body of this report, and relevant publications are listed in the References section. Section 2 describes the analysis pursuant to estimating clear-sky atmospheric transmittance from satellite data. Section 3 summarizes the research pertaining to remote sensing applications for boundary layer clouds.

2.0 SATELLITE OBSERVATIONS OF CLEAR-SKY RADIANCE

The improvement of the Air Force Geophysical Lab (AFGL) atmospheric transmittance/radiance code (LOWTRAN 6) to include a maritime aerosol model and scattered solar radiation provides the opportunity to intercompare the modeled scene radiance with satellite image data. The objectives of these comparisons are: (1) to evaluate the sensitivity of the LOWTRAN 6 code to local haze and aerosol conditions, (2) to check the correspondence between calibrated satellite radiance and model estimates, and (3) to identify means of enhancing the information content of satellite image data.

2.1 <u>Sensitivity of LOWTRAN 6 to Boundary Layer Conditions</u>

Surface meteorological data, sea surface temperatures and atmospheric sounding profiles available for the San Diego area during the 15-18 April 1986 experiment period were used to initialize the LOWTRAN model for shortwave and longwave radiance calculations. The sounding profiles provided by ship-launched balloons off Pt. Loma supplied marine atmospheric parameters up to 500 mb, and the NWS 12-hourly rawinsonde data were applied above this level. The April data set utilized for LOWTRAN is summarized in Table 1.

Direct evaluation of LOWTRAN 6 sensitivity to atmospheric fluctuations were approached by specifying each of the required parameters for the radiance model, and varying those for which some uncertainty was expected. The parameter ICSTL is the most uncertain, since it is a qualitative measure of the prevalence of continental aerosol in the marine boundary layer. If ICSTL is specified to be ten, the aerosol is entirely continental in the size distribution and optical characteristics. ICSTL set equal

to unity infers an entirely marine aerosol layer near the surface (0-2 km). Variations in longwave radiance and visibility calculated with LOWTRAN 6 were found to be sensitive to correct choice of the surface wind speeds and relative humidity in particular. Table 2 provides selected comparisons of these values for varied input parameters. When using observed values of these and the other meteorological inputs, it was estimated that the values for ICSTL should lie between three and five for the April study period in order to reproduce visibility conditions. However, radon measurements obtained from ship indicated that the ICSTL parameter should be scaled to unity, at least for the two-day period of April 15-16. It is difficult to reconcile the calculated visibility under the measured values for ICSTL, surface winds and surface relative humidity. The relatively low visibility of the marine air would seem to require higher relative humidity to be assumed at the surface. The LOWTRAN 6 model should perhaps utilize the higher moisture contents measured above the surface.

Durkee (1986) has provided evidence from a case study of satellite channel radiances that the quantity L/E, where L is a shortwave upward-directed radiance and E is downwelling irradiance, varies with wavelength in a manner that is related to the aerosol population. This quantity is equal to the directional reflectance factor for hemispheric incidence, if multiplied by the constant π steradians. We shall refer to this quantity as directional reflectance (DR $_{\lambda}$), with the assumption that it is specific to a narrow wavelength region centered on the wavelength λ . Durkee (1986) related the ratio of DR $_{\lambda}$ values in the red (λ = 0.63 microns) and near-infrared (λ = 0.86 microns) satellite channels of the NDAA-7 AVHRR and NIMBUS-7 Coastal Zone Color Scanner (CZCS) to the presence of continental aerosol. We have utilized the atmospheric conditions during the April

field program to calculate radiance in similar wavebands. For later comparison to satellite data collected in April (see Section 2.2). Channel 1 (red) and Channel 2 (near-infrared) wavebands were chosen to match those of the NOAA-6 and the NOAA-9 AVHRR instruments, which are quite close to the NDAA-7 AVHRR channels. The results of the LOWTRAN model calculations in these wavebands are summaried in Table 3 showing the dependence on aerosol characteristics, including aerosol origin, concentration and degree of hydration. Specifying two atmospheric paths in which the 0-2 km boundary layers have equal transmittance at $\lambda = 0.55 \, \mu \text{m}$ (equal visibility), but different aerosol populations, the atmosphere with a marine aerosol layer has a notably lower near-infrared (NIR) radiance, while the red-waveband radiance is only slightly less than in the case of a rural aerosol population. For the afternoon conditions observed on April 18 during the field experiment period, and assuming marine aerosol and a 23 km visibility, NIR radiance is $0.29 \text{ mWcm}^{-2}\text{sr}^{-1}$ while red radiance is $0.42 \text{ mWcm}^{-2}\text{sr}^{-1}$, and the ratio of radiances for red vs. NIR is 1.45 for the marine aerosol, while it is 1.62 for the assumption of a rural aerosol composition. Relative humidity effects are manifest through absorption of near-infrared radiation in the presence of any aerosol, and through shortwave extinction by haze droplets in the presence of the deliquescent marine aerosol. For example, a shift in the surface relative humidity from 88% to 68% for an entirely rural aerosol population decreases the red/NIR radiance ratio, primarily by increasing NIR transmittance. For comparisons such as these, the radiance ratio is sufficient, since the second multiplicative term in the expression for the DR1 ratio is the NIR/red downward irradiance ratio, which will be a constant. Quantitative comparison of DR_{λ} ratios in the satellite channels is accomplished below.

2.2 Correspondence of Satellite and LOWTRAN 6 Radiance

NDAA-6 and NDAA-9 satellite data from the Advanced Very High Resolution Radiometer (AVHRR) imager were archived during the period 15-18 April (see Table 4) from which the upwelling radiances observed at satellite position were determined. Subsets of these data are provided for the study region in Figures 1-5. These radiances were then compared to results of LOWTRAN 6 model calculations, by incorporating the boundary layer and free atmosphere contributions, as described in Section 2.1, for times matching the satellite overpass.

The satellite-observed radiance has been modeled for the viewing geometry of the polar-orbiting satellites determined from orbit parameters provided by the Scripps Satellite Oceanography Facility (SSOF). LOWTRAN 6 results were compared to satellite data by two procedures. First, spectral radiance from LOWTRAN at a specific central wavelength was matched against the equivalent-blackbody radiance of the satellite thermal channels. The central wavelength was chosen to be that at which the satellite channel calibrations had been performed, for the observed temperature range. AVHRR channel calibration is determined for monochromatic wavelengths only, and this is a central wavelength rather than the peak transmittance wavelength. For example, peak transmittance for NOAA-9 AVHRR Channel 4 is found at 10.50 microns, while radiance calibration is valid at 10.76 microns, near the center of the filter bandpass.

A second procedure was utilized for analysis of shortwave radiance, based on the channel calibration with respect to scene albedo over the entire bandpass. The LOWTRAN 6 filter function subprogram, LOWFIL, provided the means to estimate albedo.

To quantitatively test the observed and calculated radiances, the signal conversion procedure specified for AVHRR shortwave channels must be followed. A pre-launch laboratory calibration of the Channel 1 and Channel 2 sensors provides the following conversion for channel radiance, $L_{f\lambda}$;

$$\mathsf{L}_{\mathsf{f}\lambda} = \underbrace{\mathsf{A}_{\lambda} \cdot \mathsf{E}_{\mathsf{f}\lambda}}_{\pi} \,, \tag{1}$$

where the albedo (A) of the satellite channel centered at wavelength λ is given by a linear relationship to pixel brightness, and Efi is the sensorweighted solar irradiance on a horizontal surface at the top of the atmosphere. E_{fl} was obtained by integrating solar irradiance with the channel filter functions (transmittance curves) in narrow wavelength intervals across the filter bands (0.53 - 0.80 μ m for Channel 1, 0.69-1.125 μ m for Channel 2). For the LOWTRAN radiance 0.15 mWcm⁻²sr⁻¹. Channel 2 pixel albedo should be 2.4%. Satellite-observed albedos at the location and viewing geometry for which the model calculations were performed (on the morning of April 18, directly off Pt. Loma in the area shown by Figure 6) equal only 1.7%. The satellite image albedos are also less than predicted values for Channel 1, and at other image times deviations of up to 1% also are noticed. Table 5 lists Channel 1 and 2 albedos at the study location observed by NDAA-9 and calculated by LOWTRAN. Images corresponding to the three individual cases are presented in Figures 2d, 7 and 8. For the April 16 case, observed satellite radiances for Channel 2 lie in the visibility range where aerosol type (marine vs. continental) has only a small impact on extinction. The location of the value of observed radiance would lead to a visibility estimate from the plotted curves similar to that which was recorded. Again, however, the data used to plot these curves would require

a high relative humidity to produce this value. Sensitivity of model results to ocean surface albedo is not sufficient to account for the difference. Although the source of the deviation could be due to instrument response or model error, it is small in magnitude. The values of the bispectral brightness <u>ratios</u> are consistent between model results and satellite data. The ratio of scaled radiances in Channels 1 and 2 can be evaluated from the satellite data as the ratio of albedos in Channels 1 and 2. Following Eqn. 1, the value of $A_{\lambda 1}/A_{\lambda 2}$ would be expected to increase with the predominance of continental aerosol for a given value of optical depth. The ratio $(DR_{\lambda 1}/DR_{\lambda 2})$ for the NDAA-6 April 18 morning satellite pass time case is 1.30 from satellite data alone and 1.37 from model calculations, showing agreement within 1% for the ratio value.

Calculations of narrow-band radiance in the longwave channels also demonstrate agreement between LOWTRAN 6 and calibrated AVHRR data. For each of the satellite chanels, the digital pixel brightness is directly related to an equivalent-blackbody temperature. The radiance associated with this value is recovered simply by using the Planck blackbody function. Pixel radiance values may be expected to equal upwelling radiance computed from the transmission model, assuming the viewing geometry and meteorological parameters are accurately specified. The NDAA-9 Channel 4 calibrated radiance at the Pt. Loma Waypoint #1 on April 16 (satellite image is enhanced to emphasize haze in Figure 4d) was 0.79 mWcm⁻²sr⁻¹, while LOWTRAN 6 calculations for the same waveband produced approximately 0.74 mWcm⁻²sr⁻¹ (see Table 2). Variations in the viewing geometry within the area, from Waypoint #1 to #9 (a distance of 0.25 degrees longitude) amount to only 0.1% of the calculated radiance. Another case study, on the afternoon of April 17, again showed NDAA-9 Channel 4 radiance (0.80 mWcm⁻²sr⁻¹)

within 4% of model radiance (0.77 mWcm⁻²sr⁻¹). Enhanced imagery for this case is displayed in Figure 10. Results for Channels 3 (3.7 μ m) and 5 (11.8 μ m) are given in Tables 6 and 7.

It was determined from the results of many LOWTRAN output comparisons that the incorporation of accurate sounding profiles above 500 mb was not significant to the accuracy of these case studies. Estimation of surface relative humidity and surface winds are critical in predicting the concentration of haze droplets. The uncertainty in sea surface temperature is unimportant, as discussed below.

The parameterization recommended by NDAA (Kidwell, 1985) for daytime determination of sea surface temperature (SST) from the AVHRR digital image data utilizes the Channel 4 (10.76 μ m center wavelength) and Channel 5 (11.85 μ m center wavelength). For example, using NDAA-9 AVHRR data,

SST (
$$^{\circ}$$
C) = 3.6569 T₄ - 2.6705 T₅ - 268.92, (2)

where T_* is the calibrated equivalent-blackbody temperature (in degrees Kelvin) for Channel 4, and T_5 is that for Channel 5. The difference in brightness temperatures between these two channels is a measure of atmospheric moisture. T_5 is less than T_* under normal circumstances. From SST measured at Pt. Loma, observations during the period 15-17 April ranged 62°F (17°C) to 68°F (20°C). Upwelling radiance defined merely by equivalent-blackbody temperatures varies by a few percent across this range, but radiance leaving the atmosphere from an ocean surface with this temperature variation is relatively constant. Comparing the satellitederived SST for NOAA-9 pass times on April 17 (Figure 10) and April 18 (Figure 11), a variation of only 0.6° C is noted. At the same conditions, LOWTRAN estimates show Channel 4 radiances to vary from 0.77 to

0.87 mWcm⁻²sr⁻¹. This is assumed to caused by the 24-hour change in boundary layer haze conditions due to modification of the air mass through windspeed fluctuations, relative humidity changes, or alteration of the specified aerosol origin. The observed or probable SST fluctuations would not produce changes of this magnitude in the satellite channel radiances.

2.3 <u>Image Analysis</u>

Modifications to the marine haze layer are manifest in the results of LOWTRAN 6 radiance calculations at visible, near-infrared and thermal wavelengths simultaneously, through changes in the size-dependent number concentrations and extinction properties of aerosol.

The comparisons of satellite radiances in the shortwave and longwave channels to LOWTRAN calculations that were described in the previous section show that the model provides sufficient accuracy and sensitivity to study image patterns produced by haze variations. The LOWTRAN results also display smooth variation with zenith angle toward the horizon (except very near 0° elevation), as shown in the solid lines of Figure 12.

Unfortunately, narrrow—band radiometric observations taken simultaneously by the Spectral Radiometer located at Pt. Loma were too noisy (vertical bars) to draw any definite conclusions on a correspondence to model results. The interrelationship of horizontal path transmittance to upwelling radiance variations should be explored further. Time—series studies of enhanced GOES and NOAA satellite data from the field projects reveal areas of haze for which the horizontal transmittance could be measured from a ship—based or land—based filter radiometer.

3.0 SATELLITE OBSERVATIONS OF MARINE BOUNDARY LAYER CLOUDS

Research under this topic has been described in the two papers previously submitted (Wetzel and Vonder Haar, 1986a, 1986b), and will be summarized below. It is evident from routine satellite imagery that marine stratus cloud is maintained over large regions by synoptic-scale forcing, but that convective motions on the cloud scale are also necessary. Local energy budgets are therefore a critical factor in governing cloud development and dissipation. The growth of haze droplets due to radiative cooling may lead to cloud formation, and the absorption of solar energy by cloud contributes to the net energy balance in rising parcels at cloud top. The size of droplets with these cloudy parcels determines both their radiative and microphysical interactions with the environment.

A technique proposed for determination of cloud droplet size parameters, using satellite remote sensing, employs the shortwave bands where water vapor absorption is small. The results of multiple scattering radiative transfer calculations demonstrate that directional reflectance in the NIR vapor window at 1.6 µm, for example, is reduced measurably when the effective radius of the cloud layer droplet population is increased. The effective radius is a ratio of volume-weighted to cross-section-weighted number concentrations for cloud droplets, which provides a parameter that scales the relationship of droplet absorption versus scattering processes.

It is suggested that a satellite sensor designed for this application should provide NIR window radiances simultaneously in at least two channels, with one near the visible (0.85 μ m) to estimate the scaled optical depth of the stratus layer. Figure 13 shows the direct relationship of scaled optical depth to 0.85 μ m reflectance, while Figure 14 is an example

of how 0.85 µm reflectance can then be used to discriminate the 1.6 µm reflectance variations due to droplet size. Several tests have been carried out with the multiple scattering model to simulate remote sensing of multi-layer clouds. The optical depth, droplet characteristics near cloud top, and viewing geometry all are significant factors of cloud reflectance, but can be separated through numerical experiment. We have also designed and field-tested a remote sensing system to measure spectral radiance in the NIR windows centered at 0.85, 1.6, and 2.2 micron wavelengths. The schematic design of the Discrete Filter Wheel implemented for this study is shown in Figure 15. The Spectral Radiometer System was operated as part of cloud microphysics-radiation experiments in collaboration with NOSC during April and June, 1986. Data from the June aircraft flights show the cloudscale fluctuations in upward-directed spectral radiance expected in the shortwave bands. Analysis of the data sets for reflectance signatures that are related to microphysical conditions in ongoing. An example of the raw data captured during the research flights in June is presented in Figure 16, where the vertical scale is proportional to measured radiance in each channel, and the horizontal scale represents fractions of seconds along an aircraft flight leg. A complete study of the results from the summer experiment will include combined evaluation of the satellite, radiometer and microphysical data sets.

Digital imagery available from a DMSP Special Sensor C (SSC) experiment during a short period in 1979 provides the only satellite observations of clouds in a near-infrared window band. Figure 17 is an example set of images in the visible, near-infrared (SSC), and thermal window for an ocean region off the East Coast. The ability of multi-spectral techniques to discriminate low water cloud from ice cloud is obvious by comparing the

brightness of the three images for the high frontal cloud band in the center of the image. The value of the Special Sensor image for studying the stratocumulus cloud field near the top of the image is not obvious. While the radiative transfer calculations indicate cloud reflectance should be lower for the SSC than the visible channel, the effects due to coarse resolution and broken cloud fields are difficult to assess. Remote sensing studies from aircraft should be continued for the stratus applications, particularly because these simultaneously provide the aerosol/droplet sampling, as well as the radiometric observations at higher resolution than satellites currently allow.

4.0 SUMMARY AND RECOMMENDATIONS

The research project during the past year has allowed us to extend th results of previous studies on satellite applications, particularly with regard to intercomparison with standard atmospheric transmittance models. It has also been possible to study the basis of a new technique for characterizing marine layer clouds, so that their evolution may be related to the near-cloud environmental parameters. A wider range of meteorological scenarios must be routinely sampled with the radiometric instrumentation available and coincident satellite-based imagers to allow improvement of and reliance on the numerical models during image interpretation. The multiple-scattering radiative transfer methods can then be meshed with LOWTRAN to provide cloudy-clear air analyses.

5.0 ACKNOWLEDGEMENTS

The successful completion of the aircraft field program described in Section 3 was only possible through the expertise contributed by Dr. Doug Jenson of NOSC, to whom the authors are indebted. Appreciation is extended to Dr. Herb Hughes of NOSC, Dr. Phil Durkee of the Naval Postgraduate School and the Gibbs Flying Service for their support. Other contributions were made by Hung-Chi Kuo, Jan Behunek and Chris Johnson-Pasqua of Colorado State University. Discussions with Dr. Steve Cox of CSU on various research aspects are also gratefully acknowledged.

6.0 REFERENCES

- Durkee, P.A., 1986: Aerosol characterization with dual-wavelength radiance measurements. Preprints, AMS Second Conference on Satellite Meterology/Remote Sensing and Applications, 13-16 May 1986, Williamsburg, VA, 298-302.
- Kidwell, K. B., 1985: NOAA Polar Orbiter Data Users Guide. National Oceanic and Atmsopheric Administration, Washington, DC.
- Kneizys, F. X., E. P. Shettle, W. O. Gallergy, J. H. Chetwynd, Jr., L. W. Abreu, J. E. A. Selby, S. A. Clough and R. W. Fenn, 1983: Atmospheric transmittance/radiance: Computer code LOWTRAN 6. AFGL-TR-83-0187, Hanscom AFB, MA, 200 pp.
- Vonder Haar, T. H., 1985: Progress Report #1 to Naval Ocean Systems Center, Contract N66001-86-R-0202. METSAT, Inc., 1 p.
- Vonder Haar, T. H., 1986a: Progress Report #2 to Naval Ocean Systems Center, Contract N66001-86-R-0202. METSAT, Inc., 4 pp.
- Vonder Haar, T. H., 1986b: Progress Report #3 to Naval Ocean Systems Center, Contract N66001-86-R-0202. METSAT, Inc., 1 p.
- Wetzel, M. and T. H. Vonder Haar, 1986a: The impact of stratocumulus microphysical variations on near-infrared radiance. Extended abstracts, AMS Sixth Conference on Atmsopheric Radiatiohn, 12-16 May 1986. Williamsburg, VA, 153-156.
- Wetzel, M. and T. H. Vonder Haar, 1986b: Modeling and observation of near-IR cloud signatures in the marine boundary layer. DoD Fourth Tri-Service Cloud Modeling Workshop, 3-4 June 1986, Janscom AFB, MA 50-64.

TABLES

TABLE 1. Data Set for LOWTRAN Model Calculations

Data Source	<u>Description</u>	Date/1	ime(GMT)
NOSC	Vertical profiles from ship-launched balloons, near Pt. Loma	16 Apr	86/0345 /1645 /2045
		17 Apr	86/0045
NWS	12-hourly rawinsonde profiles from Montgomery Field		86/1200- 86/0000
NOSC	Pt. Loma surface observations (inc. sea surface temperature)		86/2230- 86/0510
			86/1530- 86/1700
		•	86/2000- 86/2100
			86/2345- 86/0130
NWS	Hourly surface observations at San Diego, Lindbergh Field, San Nicholas Island and North Island	•	86/1200- 86/0200

TABLE 2. LOWTRAN Model Parameters and Results for the 10-11 üm band.

Time (GMT)	Surface Relative Humidity (%)	ICSTL (n.d.)	Current Wind Speed (m/s)	24-hour Average Wind Speed (m/s)	Visibility (km)	10.76 üm Radiance (mWcm ⁻² sr ⁻¹ ųm ⁻¹)
						•
0045	* 50	1	10.3	7.7	32.7	0.75
(17 April)	60	1	10.3	7.7	27.0	0.74 .
	60	1	9.8	7.7	28. 1	0.75
	60	1	7.7	7.7	33.4	0.76
	60	1	1.7	7.7	66.0	0.80
	60	1	10.3	5.1	31.9	0.74
	60	2	2.6	7.7	57.3	0.80
	60	3	10.3	7.7	24.8	0.74
	60	5	9.8	7.7	21.8	0.75
	65	2	7.7	7.7	29.3	0.76
	70	1	7.7	7.7	28.2	0.76
	70	2	7.7	7.7	26.8	0.76
	75	2	7.7	7.7	24.8	0.75
	98	1	10.3	7.7	2.9	0.47
2045	* 50	2	10.3	6.2	34.6	0.74
(16 April)	65	1	10.3	6.2	27.1	0.73
	65	2	10.3	6.2	26.1	0.73
	75	2	7.7	7.7	24.8	0.75
1645	* 65	1	4.6	6.2	50.1	0.77
(16 April)	75	1	4.6	6.2	43.1	0,77
	85	1	4.6	6.2	38.8	0.77

Notes: Upper air temperature and humidity data obtained from radiosonde ascents for 16 April 1986. Parameter lists marked by an asterisk(*) represent observed surface conditions of relative humidity, ICSTL, and winds for the time period indicated. Calculated visibilities are greater than those measured at 1645 GMT (18.5 km), 2045 (27.7 km), and 0045 (18.5 km).

TABLE 3. LOWTRAN Model Parameters and Results for NOAA-9 AVHRR Channels 1 and 2

Time (GMT)	Channel Number	Surface Relative Humidity (%)	ICSTL (n.d.)	Current Wind Speed (m/s)	24-hour Average Wind Speed (m/s)	Visibility (km)	Filter-Integrated Radiance (mWcm ⁻² sr ⁻¹
0045	1	60	1	10.3	7.7	27.1	0.163
(17 April)	1	60	2	7.7	7.7	32.0	0.158
	1	60	5	10.3	7.7	21.2	0.171
	1	75	2	7.7	7.7	24.8	0.166
	2	60	1	1.7	7.7	66.0	0.104
	2	60	3	10.3	7.7	24.8	0.141
	2	60	5	10.3	7.7	21.2	0.147
	2	60	5	7.7	7.7	24.9	0.137
	2	60	5	1.7	7.7	39.5	0.113
	2	70	1	10.3	7.7	26.8	0.139
	2	70	1	7.7	7.7	28.1	0.137
	2	70	1	7.7	10.3	23.7	0.145
	2	70	1	5.1	10.3	29.7	0.134
	2	70	5	10.3	5.1	19.5	0.150
	2	75	2	7.7	7.7	24.8	0.142
	2	80	5	10.3	7.7	14.4	0.170
	2	95	1	10.3	7.7	10.1	0.2 02
	2	95	5	10.3	7.7	6.3	0.215
	2	98	1	10.3	7.7	2.9	0.239
	2	9 8	5	10.3	7.7	1.5	0.247

TABLE 4. NOAA AVHRR Data Archived For April Experiment

<u>Satellite</u>	<pre>Date/Time(GMT)</pre>				
NOAA-9	April 15/2227				
NOAA-6	April 16/0246				
NOAA-9	April 16/2216				
NDAA-6	April 17/1442				
NDAA-9	April 17/2205				
NDAA-6	April 18/1418				
NOAA-9	April 18/2155				

TABLE 5. LOWTRAN model parameters and results for NOAA-9 AVHRR Channels 1 and 2 $\,$

Time (GMT)	Channel Number	Surface Relative Humidity (%)	ICSTL (n.d.)	LOWTRAN Albedo (%)	Satellite Albedo (%)
2227 (15 April)	1 2	70 70	1	3.2 2.0	3.0 1.0
2216 (16 April)	1 2	75 75	2 2	2.9 1.9	3.4 2.1
2205 (17 April)	1 2	50 50	1	2.0 1.3	3.2 1.7

TABLE 6. LOWTRAN Model Parameters and Results for NOAA-9 AVHRR Channel 3.

Time (GMT)	Surface Relative Humidity (%)	ICSTL (n.d.)	Current Wind Speed (m/s)	24-hour Average Wind Speed (m/s)	Visibility (km)	3.73-üm Radiance (mWcm ⁻² sr ⁻¹ um ⁻¹)
0045	60	1	10.3	7.7	27.0	0.027 .
(17 April)	60	1	5.1	7.7	43.5	0.028
	60	5	10.3	7.7	21.2	0.027
	80	1	10.3	7.7	19.7	0.026
	95	1	10.3	7.7	10.1	0.022

Note: Satellite radiance for 2216 GMT (16 April) pass is 0.026 mWcm $^{-2}$ sr $^{-1}$ um $^{-1}$ at Way Point #1 near Pt. Loma.

TABLE 7. LOWTRAN Model Parameters and Results for NDAA-9
AVHRR Channel 5.

Time (GMT)	Surface Relative Humidity (%)	ICSTL (n.d.)	Current Wind Speed (m/s)	24-hour Average Wind Speed (m/s)	Visibility (km)	ll.8-üm Radiance (mWcm ⁻² sr ⁻¹ um ⁻¹)
0045	50	2	5.1	7.7	49.7	0.75
(17 April)	60	1	10.3	7.7	27.0	0.72
	60	5	10.3	7.7	21.2	0.71
	80	1	10.3	7.7	19.7	0.70

Note: Satellite radiance for 2216 GMT (16 April) pass is 0.72 mWcm $^{-2}$ sr $^{-1}$ um $^{-1}$ at Way Point #1 near Pt. Loma.

FIGURES

shown.

(in

San Clemente Island

coastline (right edge of Fig. 1(c)). Channel 1 data

between

areas

data in three adjacent

San Diego

of panel (a), above, and

satellite image

1. NOAA-9

•

Island

Clemente

San

central

_

Channe 1

1(b).

Fig.

30

Fig.

끐

9

APRIL

+ 1.50

0.292)

value *

(pixel

Albedo(2)

~:

Channe 1

まままままままらていいじゅんりんかんかん にんこうこう こうこう ちょうきょうきょうしょう こうこうしょ しょっしょうしょうしょうしょうしょうしょうしょう にんかい かんかんかん こうしょうしょうしょうしょう とうしょう きょうしょ こうこうしょ こうしょう かいかい かいかい カート・ション こうしょう しょうこうしょう すす すり うちょうごと とこご なんじょ ととご ごご と なん なんか かんかんご となること なら こう うりょう マラリ しょう しょう しょう しょう しょう しょう しょう しゅうり しゅう **!!!!こごごごにならならならなごごり!!ごごららめいかんかいんなごごごごとなるしょうこごとすることです! !!!!!」ここことにににににこことすここことにいたサレナリシににごごごとなるこう!~!!!!!!!!!!!!!** てままままし くこう にんこう たんしょ こうしょう しゅうしょ トルトトトトラー こうしょう こうしょう こうしょう しょうしょう ままままたこうことにはなるなどにもなりのとかりからにはなりというないとことととなっていることもままましょう すすすすすす ことごにににままなどごのどはなめにはなわかかもなまなことはないとなるですです。 11112 とまたをとなってまるものもとなわかかかかならららりとくととなるのもくとなるとなってもとととととなっているととなるとなっているというというというというというというというというというというという これではなられるなくないないない。 **各型型中有中央等等的有中央等等等等等等等等的等等等的的现在式息—————** € ((

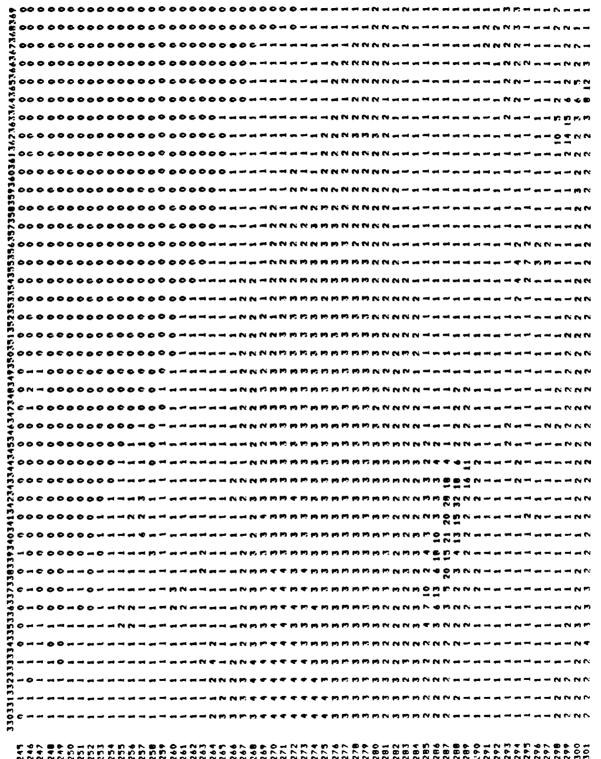


Fig. 2(b). Channel 2, central panel of image area.

(

~

3

Fig. 2(c). Channel 2, eastern panel.

2

APRIL

Fig. 2(d) Enhanced satellite image, showing area of three adjacent panels in the Fig. 2(a)-2(c) in center, and surrounding region.

Channel 3: Equivalent-blackbody temperature (deg Celsius) = (pixel value * 0.245) -13.74

crons	(•	(((C	(•	C	C	:	?	•	Э	•	7	3))))
Bandpass is 3.44 - 4.03 ml	A-V APRIL. 15 72152 CH. 3	29029129229396297298299300301302303304305306307308309310312313314315316315318319318312313233233243293263273283 1721221221221221221221221212122122123124124123121122121241291231231231231231231231231211211191211221231221231231231231231231231231231					11811911511413016516416113712212011712012112112112112011911911811818181201201211201211212222121122121212012012			11912012012012011811181181181161361711661641701681611551132011911811811811911911911911201301191201201201201201201301191191191191181 1191201201201201181161171181331711711701731741751641531221221191191191191191191191181181181181201201201201191191191181181				119117119120171191201201191201191201201201201211211201211211221231231231201211221231261261261261261241241241241281281281281281281171171191171171191191171171201212121212121212121212121212121			20120 7118 8 7119 8 20 8 8 9 9 9 9 9 9 9 20 20 20 23 23 23 23 23 34 30 23 23 23 23 23 23 23 23 23	1221201201201201191181191191121121191191191191191191191191191		1181191191191201191191191181181181181181161 12012112012212011911911912011911911911811 1201201201201191201181181181181191181161	(a) Ac in Fig. 1 for Channel 3 data
	((٠	•	((· ·		Ć	C		,	ر	,	j	J	,)	Ć		j

Fig. 3(a) As in Fig. 1, for Channel 3 data.

)

)

)

Fig. 3(b). Channel 3, central panel.

7

)

4

ì

)

)

(

3

NPR II.

--

(

(

(

(

(

Ĺ

(

(

37

0

)

)

,	,
:	
4	
•	-
1000	-
2	È

しゅうてんききゅうそう ほうてくんゅうきゅう ちゅうきゅうりゅう ほうりょう こうこう こうしょう しょうしょう しょうてい しゅうこう とり しゅう	55
そですててまじまり こじょりだだり ロングンをジョングルット トット・マット・マット・アット・アット・アット・アット・アンジン イン・アンジン かんしゅんこう イン・アンジン くり シャン・アンジン くり シャン・アン・アンジン くり シャン・アン・アンジン くり シャン・アンジン くり シャン・アンジン くり シャン・アンジン くり シャン・アンジン くり シャン・アンジン くり シャン・アンジン ストー・アー・アンジン ストー・アー・アンジン ストー・アー・アンジン ストー・アー・アンジン ストー・アー・アー・アンジン ストー・アー・アー・アー・アー・アー・アー・アー・アー・アー・アー・アー・アー・アー	F -
ARDER BERTER BERTE	75
in Filler in France and Cooke and Co	611
	1 1 6
ノノイ シェンティイン との とくごに かりふめる イザ くららく こくく とくしょう こうしょう こうしょう しょくしょく しょくしょく しょくしょく しょく しょく しょく しょく しょ	10
- トローロット・ロット・ロット・ロット・ロット・ロット・ロット・ロット・ロット・ロック・ロック・ロック・ロック・ロック・ロック・ロック・ロック・ロック・ロック	181
「ドリュアリアにしょっていませましょう。 アンジャン・アンジューマー・アンジャン・アンジャン・サンシャル・サンシャル・サンシャン・アンジャン・アンシット	75
	911
	:==
982222222222222222222222222222222222222	:53
	5.6
648876498164446000000000000000000000000000000000	5.5
	2.5
опетация по	12
	7 - 6
	23
のおおんり () そんしょく () かんしょう () かんしょんしょんしょんしょんしょんしょんしょんしょんしょんしょんしょんしょんしょん	25
- V トレブリモ ドビンジャードレーン C クロンター・ジョー・ジョー・ジョー・ジョー・ジョー・ジョー・ジョー・ジョー・ジョー・ジョ	131
	77
	913
	22
	119
940 40 8 8 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9	66
	201
	5.5
	911
	3 2 2
	118
	119
2888798798900800000000000000000000000000	1202
0.00 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	200
	6
	011
	11
	===
	200
	181
	911
	120
)	25.5
	-0-
M	
	900
44444888888888888888888888888888888888	i be: be)

C

Fig. 3(d), Enhanced satellite image for Channel 3, showing area of Fig. 3(a)-3(c) and surrounding region.

(H-9 APRIL	14.22 INC. 13.10.10.10.10.10.10.10.10.10.10.10.10.10.
(29029129 45 17817919 46 17717817	22932942952962972982993003013023033043053063073083093103113123133143153163173183193203713233334325373283732837 01901801801801811811811811811811871821821821821831331831831821821821821821821821821821821821821821
(71	61/7/81/H101/71/71/918018018019118208218218218218218218218218218218218218218
.(50 16416516 51 18017517 52 18418718	6916916917117217617717717818801801801801818118117918017938017938018018018018018018181181181181181181181
(416217 416519 016818	80.45714467167167167167167167177177177177777777
•	57 16416919 58 16616116 59 16215116	88 32 72 62 62 63 63 65 66 57 67 53 68 64 70 70 71 72 74 75 75 75 77 77 77 78 78 78 78 80 79 79 80 8 8 8 74 86 80 75 66 64 65 66 67 67 67 66 66 67 67 67 70 72 73 75 75 77 75 75 77 77 75 76 77 78 79 79 79 79 79 79 79 79 79 79 79 79 79
(.	61 16316516 62 16316516 63 16316316	
(64 16416316	016116016216718719318618718517417617117017017017217217317417417417417317317317317317317317317417417417417417417 91591411671621801951911881881811771771701701701701711711731731731731731731731731731731731
æ	161161162	62162162162180184187187181181181179173168169169179168170170168168168167169174177177170170170170170176516516716 621631631631631731641881891891891891891717117017617016916616716716716716717117217016316416516216316316416516716 63163162162163163173191188189189189189187818317817116781671671641651651651661661661651551631631631631631631631
C	1631631	64163163163161317711711711717170719719719719617317216516316416516316316316316316716716716716716716716716716716 641631631631631631611771171171193197197197196173172165163164716316516616516616416416716716716716716716416416416 661671631621611621671671671919019719819719619719619017416716516516316316416516516516416416716116016116316416416
¢	16516616 15916915 16916916	671651621611611641631611631761941991971991981981R01R91R716916716616316416516516516616416165161161163164165167 67166164164164164163163163163167172189196186174169170173170168166166167167166165168166167165164164164164164164 6616616616616616616516416416416316717317017016916917017016R167169167169168168168168166165165164163165165165
•	4816 6716 4516	6416516516416416416516516516316216316416916716717017117217017117117217217117016716716716716416716716 6616716716716716616616616616316216316316116516716916917017317217317317317317317016916716816816916716716 691701701691671671671681671661631631631671671691691671171172173174176176175173171171171717015916916916
ī.	278 1691671681 279 1701701701 280 1691681711 281 1471481681	69171170171169168168168167167167164165167169170171116917217317317417617617717717517217317417417417717 7017017117117116916816916916616616616816818717217417417417517517517717717817817717761781777751751741741777777 7117117172172169169167163168170169172174177173174174175177177177177179178178178177177176178177777617761777777
3	1691681	7015017117117016816416417417217217217217217317417517617517747717717177177177177177177175175175175175
3	1501691	001/01/41/016/1661/11/31/31/31/31/31/11/31/31/41/61/31/31/31/31/31/31/31/31/31/31/31/31/31
ز	1711651	7017017117317317317217117017017117711721731731711721731721731721741741741741751741761771781791791791791791791 711711711711771771617116171
3	171172171	
)	174174171	7.171171171170171171170167168168168716917017017176176176176177176176176176176178178178178178178178178178178178 7.217717171701711701711701671681671691701701701761761761761777777777777
•	75174 76176 76176	74 73 72 72 73 73 73 73 75 75 75 75 77 77 77 77 77 77 77 77 77
)	0 17817177 1 178179178 1 178179178	

)

כ

)

)

)

(

1

(

4(5). Channel 4, central panel. . . .

Ĺ

(

9

1

(

(

C

(

(

(

C

C

:

3

3

)

)

7

)

)

)

)

)



Enhanced satellite image for data shown in Fig. 4(a)-4(c) and urrounding region.

,		
í	245	2×1.2×22×3×2×42×2×62×2×2×3×3×3×3×3×3×3×3×3×3×3×3×3×3×3
	240	Y LOOT / Y LYY LOOT (Y LOT LEAT LEAT LEAT LEAT LEAT LEAT LEAT LEA
(248	2174174175176176176178178178178179180180180181182182187182187187187187181181181181182187187187181181181181181 81701721741701731751771791791791801811811811871871821871818118118118118118118118118118118119181191181197187187
(250	
	253	VIOU BSIGSIVOIVOIVOIVOIVOIVOIVIISIVIVAVAIVAIVAIVAIVAIVAIVAIVIVIVIVAIVAIVA
(254	621611621651671641c51661701701691721721721721731761751751761761761771781791791791801811811811811871871 601621621651631641651661661671681701711711721731731751761771781781781791791791801811801811801811801811
(256	.41581591591631631641651651661651671681681681711731741731761761761771781771781781781781801801801801811811 *01591591591611631641661551641661671651681671717417517317317317417617717717717717717717777777777
	258	7717216416216316316516516516416316316516516516616817017217017217317317617617617617617617717171771781781791791 821791581631641641651521671631641651651691671571681691681691701711741741691691721731751761771761781781771
(260	c 59 60 70 88 84 72 65 63 62 69 67 67 68 68 64 64 64 68 68 68 67 67 72 72 73 76 76 76 77 77 77 76 76 76
	262	01901861831791791721701701701701691691671711701721711721711721711701721721721721721721721721741711701731741 31691911701941851811711671681701701721571681671671571581701721711721711711701721711771717171717
_	264	9018318418317017516916816816816817117117117217217317711711711691711711701771701701731731731731731
(266	0160177181187184181179177517216516716716816816816716716716717317731731731781681671671681671671671671671
	267 268	//185196 83 /9132 82/30 /1 68 67 67 67 67 67 65 65 65 65 65 69 /1 58 6/162 60 59 58 61 62 63 65 62169 88 85 85 85 88 86 82 74 69 65 64 64 65 65 62 62 62 63 63 63 53 56 56 59 57 56 56 56 60 60 60 60
	269	401401611401671891891891871901731841911831671671421631621601601601601601601601601601601601601601
	271	\$410815015015016710881901771771771771771771771771871871871871871
	272	31641641651621601581581621601581591771911971941961971971761881871661651641601621631621631631631631531531591581601611621621
	274	87167167167167167167167167167167167167167
)	275	67163163163163163163163163161158150162164165167157163167157169170168169169170170168168169165165165165165165165
ı	276	01681671651651651651671671671651 81691701681681671571691581661
	278	716:1661671691681691671661661651651591601631641661671681661701711721731751761761701711751731721701701701701701
	279	681.671.641.641.641.571.611.571.701.701.701.731.731.731.751.761.781.781.781.781.751.771.751.741.731.731.731.71 471.451.401.481.461.461.701.731.731.761.781.781.781.781.781.781.781.781.781.78
	281	84170171171187187187187170171871871871871771787178
ı	282	516216417317117017217017016917017317417517417417517517617617517517517517741741741741771771751741771 2158174172168173171701711701711771771781781781751751751781771771771771771771771771771771771771
	184	681711681651631701721701711711731681691711721731751731741731741721701691701721731731741731741731
1	28 5	731731711771701721731721701701711711721731741741741751741761 7317117217111721741721721701711721721781761761761761761761
	287	31721711701711681691581681731741741721701701701711721721711731741751741751771781781781781781
	28 8	11701691681671681701701701731721701681701711701701721721721731731731731741751771791791791781781781 0167169164147171177173172172177171701721711721721701721721731741741741741751771791791791791791
	290	11461681651671481711731731691711731731731731731771701701701711731741761761771761781791781791791
	295	
	2 4 3	591581691691701581701701681551551681691661681721751751751751751751751741741741741771771781771781771781771781791
	(1) (1) (4) (4) (4) (4)	91681631661651671671671701751751751751751761761761751751751761761761761761761761761761761781781781781 0169169170170171171171771741751751751761761761761771751751751761781781781781781781781781781781781781781
١	296	11711711711711741751741741751761761761761761761771761761761761761761
	297	717717717617717617617617617617617617617771 81781761761761761761761761761761761771771741
,	566	241251751261751251251251251751751751751751751751761751761781781781781771761761761761761761761761761761761761
	300	171 ** ** ** ** ** ** ** ** ** *
	;	

3

,

)

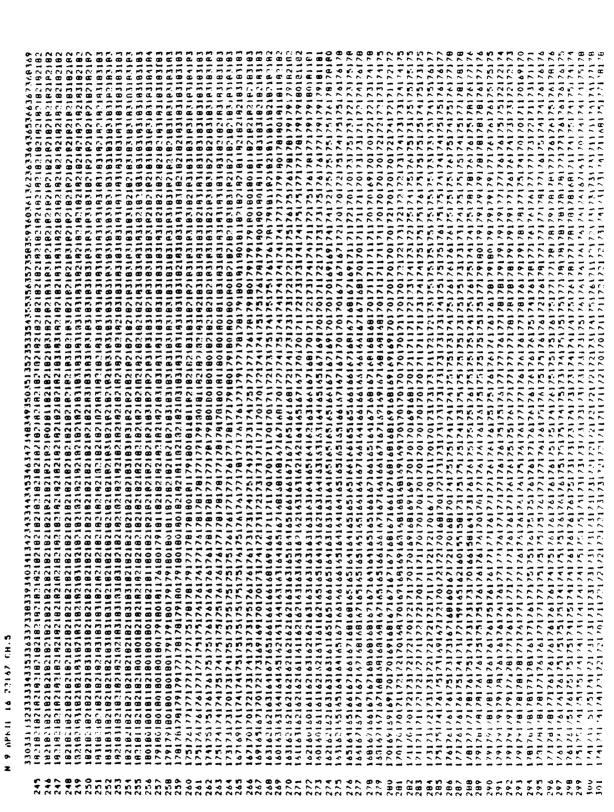
)

)

)

Fig. 5(1). A un Fig. 1, for Channel 5 data.

)



3

)

(

(

)

ţ

~~~~~~~~~~	002203 004202 004202 005206 110210 112210 119221 119220 112220 112220 112220		. ^ ^ 8 8 8 8 4 4 4 4 4 4 4 4 4 4 4 4 4 4
8 4 7 6 4 - 7 7 5 8 9 8 9 9 8 6		-40405050606066	01801111118011
4000000000000000		20000000000000000000000000000000000000	001111111111111111111111111111111111111
Z C C C C C C C C C C C C C C C C C C C			
222222222222222222222222222222222222222		0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	. ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~
90077777777000	**************************************	883 83 77 77 77	7 7 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1
888888888888888888888888888888888888888		4162 4163 4163 4163 4163 3162 3162 2162 2163 2163 4176 8176	777888888888888888888888888888888888888
	20000000000000000000000000000000000000	7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7	341444444444444444444444444444444444444
		2 4 4 4 5 6 6 6 6 6 6 7 4 4 4 4 7 7 7 7 7 7 7 7 7	771 771 771 771 801 801 801 801 801
			771 771 771 771 771 771 771 771 771 771
	> = = = = = = = = = = = = = = = = = = =		30 37 37 37 37 37 37 37 37 37 37 37 37 37
		601222223333	222222222222222222222222222222222222222
************************			21 22 23 23 23 23 23 23 23 23 23 23 23 23
			7751 7751 7751 7751 7751
*		~	51751 51751 51751 51751 51771 51771 61781 61781 11871 11871 11871
22163 2163 2163 2163 2163 2163 2163 2163		# M # # # # # # # # # # # # # # # # # #	9799777777777
			4054000000
	NAU NE CE CE CE PE PE	8831 8821 8821 8821 8821 8921	721 721 721 721 721 721 731 731 731 731
	621 621 631 631 631 631 631		791177791177791177791177779117777777777
222222222222222222222222222222222222222		833 183 183 183 183 183 183 183 183 183	1880 172 173 173 174 174 180 180
		3555555555555	3 ED 60 / / / / / / / / / E 60 ED
			20118 80118 80118 7317 7317 7517 7517 7917 7918
			7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7
	<b>9 C B B B C C B B B B B B B B B B B B B</b>		77 77 77 78 80 77 78 80 77 78 80 77 78 80 77 78 80 77 78 80 78 80 78 80 78 80 78 80 80 80 80 80 80 80 80 80 80 80 80 80
9380 2182 2182 2182 2183 3183 3183 3183 3183			217777
			8 7 8 7 8 7 8 8 8 8 8 8 8 8 8 8 8 8 8 8
		2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	7 2 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3
~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~			777
13/2/2011		0	300000000000
			27779977777
**************************************		80118 8018 79018 7718 7718 7718 7717 7917 7917	
			2
			~ <b>~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ </b>
487642200000000000000000000000000000000000	0 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4	2	9 8 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9
14 14 14 14 14 14 14 14 14 14 14 14 14 1		, , , , , , , , , , , , , , , , , , ,	

C

C

N-9 APKIL 16 22162 CH.5

_)

Fig. 6. Enhanced satellite image, showing immediate study area.

Fig. 7., Enhanced satellite image for 15 April NOAA-9 pass, showing area of, pixel sampling within box.

61 308

Fig. 8. Enhanced satellite image for 17 April NOAA-9 pass.

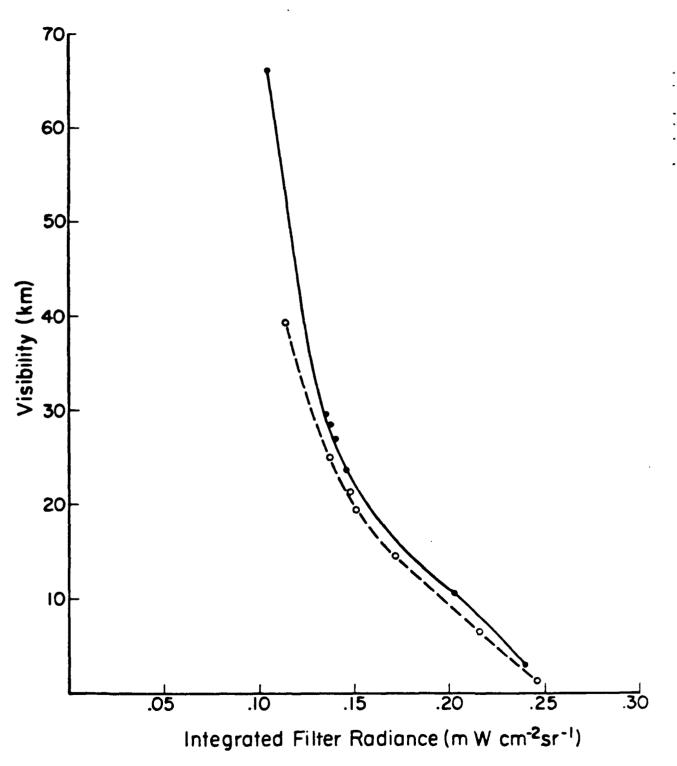


Fig. 9. Radiance calculated by LOWTRAN 6 for AVHRE Channel 2 encar-primated and resultant visibility for a range of surface meteorological robust. The Values of ICSTL parameter are 1 for solid circles, and 5 for open careful Lines represent hand-drawn curves connecting the values.

Fig. 10. Enhanced satellite image for 17 April NOAA-9 pass (Channel 4).



5.2

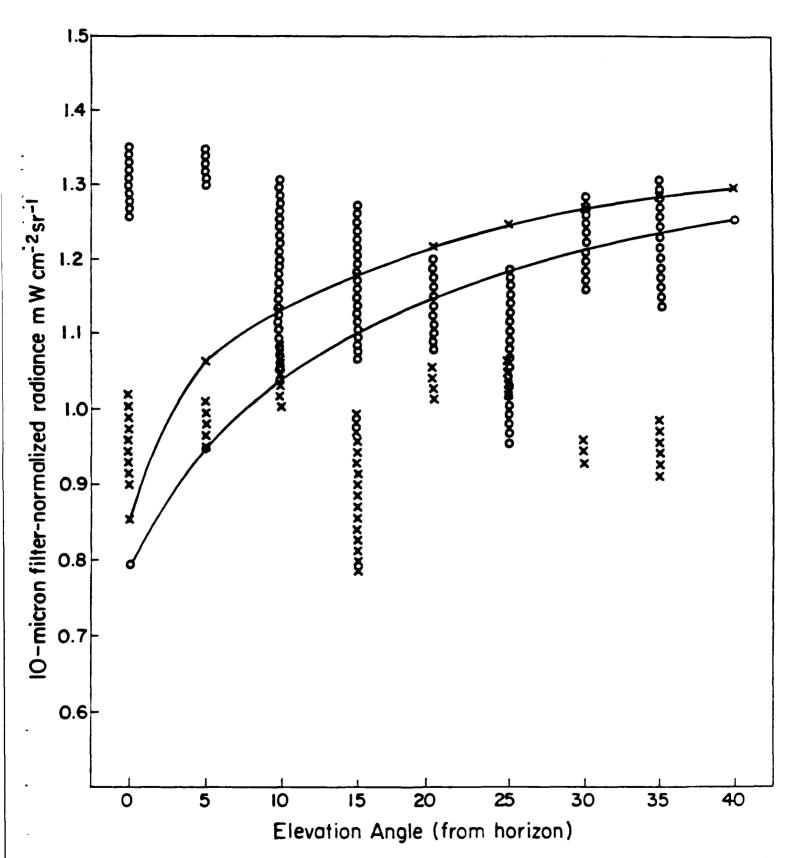


Fig. 12. LOWTRAN 6 radiance calculations for the 10-micron filter band(solid curves) used to obtain Spectral Radiometer clear air measurements (range of values shown in vertical bars of X and 0 symbols). X refers to data from 18 April; () for 16 April.

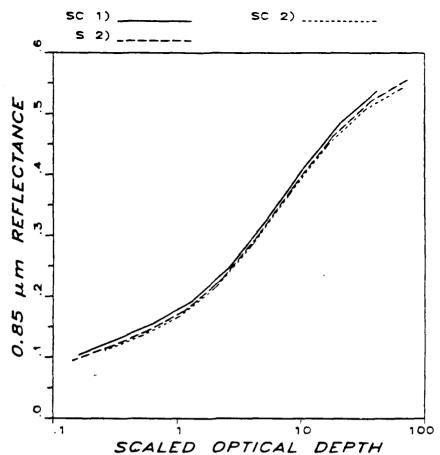


Fig. 13. Reflectance calculated for three different model cloud layers.

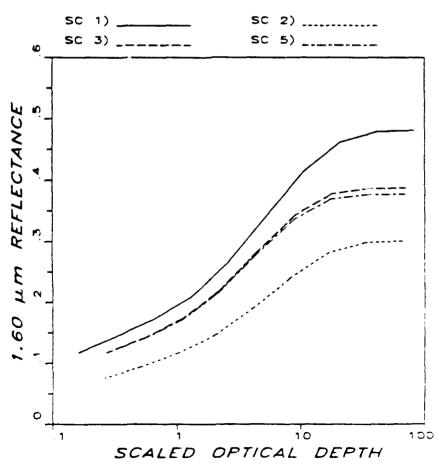


Fig. 14. 1.60-um reflectance calculated for different model cloud layers. 54

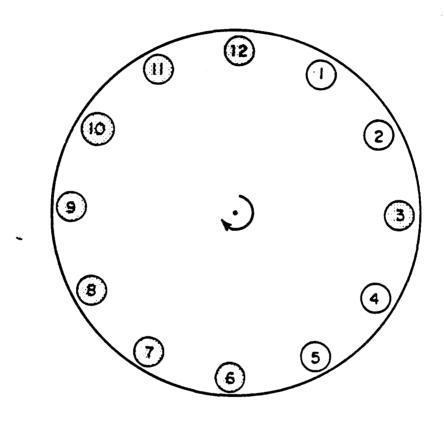
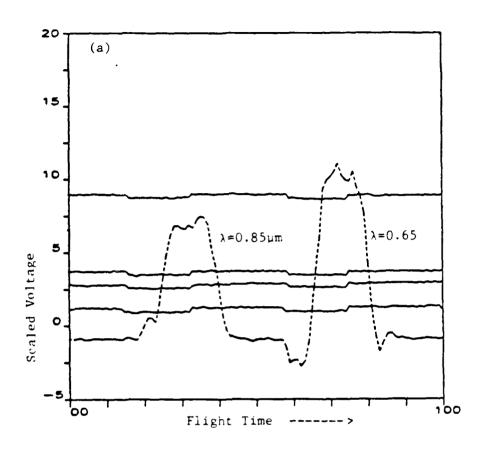


Fig. 15. Schematic diagram of Spectral Radiometer discrete filter wheel. The central wavelengths of each filter are listed below:

Filter #	Central Wavelength	(mm)
3	3.7	
6	2.2	
7	0.85	
8	0.65	
9	0.85	
10	1.6	
11	1.6	
12	10.5	
(1,2,4,5)	(open)	



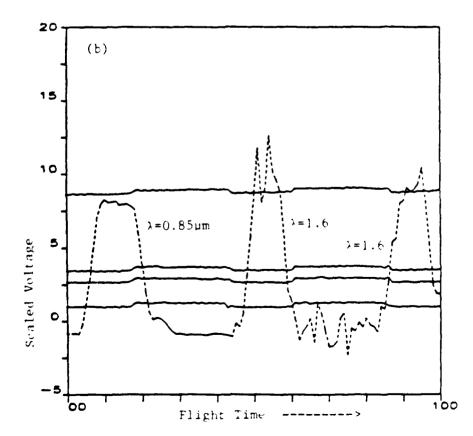


Fig. 16. Spectral radiometer (dotted) and shortwave flux (solid) data for stratus flight studies.

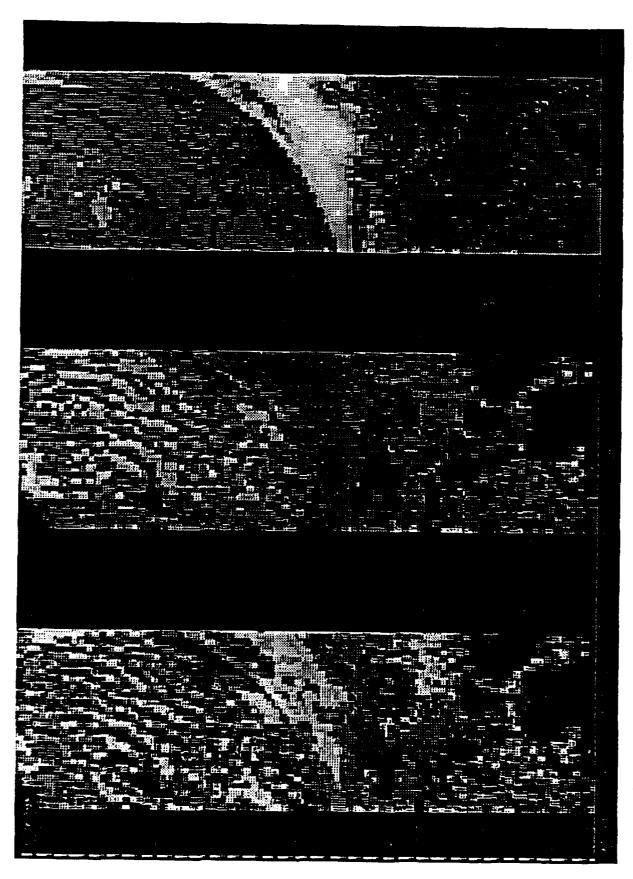


Figure 17. DMSP images in the visible (left), SSC (middle), and infrared (right) channels.

## END DATE FILMED

E.A.